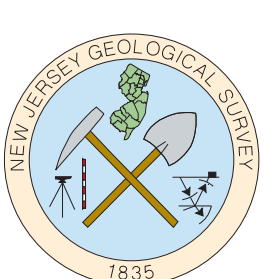


Base from U.S. Geological Survey, Roselle, N.J. 7 1/2-minute quadrangle 1955
Photorevised 1981
GIS application and digital cartography by:
Donald H. Monteverde

SCALE 1:24 000
1 MILE
0 1000 2000 3000 4000 5000 6000 7000 FEET
1 KILOMETER

Surficial geology by S.D. Stanford, 1991
Bedrock geology by D.H. Monteverde, 1999-2000
Review of these interpretations is limited to the New Jersey Geological Survey



**BEDROCK GEOLOGY OF THE ROSELLE QUADRANGLE,
UNION AND ESSEX COUNTIES, NEW JERSEY**

by
Donald H. Monteverde

2000

LOCATION IN
NEW JERSEY

Introduction

The Roselle 7.5-minute topographic quadrangle lies within the Piedmont Physiographic Province in north central New Jersey. Union County covers the majority of the quadrangle while Essex County occupies the northern third. The quadrangle is highly urbanized with large open space parcels restricted to the northeastern side of the Watchung Mountains.

Stratigraphy

Surficial cover dominates the quadrangle geology. Detailed surficial geologic interpretations have been done by Stanford (1991). The well log data (table 1, simplified from Stanford, 1991) supplemented by gravity measurements (Ghage and Hall, 1991) allowed the contouring of the bedrock surface beneath these recent sediments (Stanford, 1991). A Pleistocene aged, deep glacially scoured basins (current elevation 0) downslope through the First and Second Watchung Mountains creating Millburn Gap and connects with second basin (outlined by 50-foot elevation trends) trending ENE-WSW (Fig. 1). The Pleistocene cover almost 90% of the quadrangle (Stanford, 1991). Ridges underlain by igneous units delineating a north-northeast trend are the only areas where bedrock emerges from the surficial cover (see Fig. 1). Rare sedimentary rock exposures occur near the contact with the first basal ridge. The remaining areas remain blanketed under Pleistocene to Recent unconsolidated sediments.

Structure

Bedrock structure consists of a gentle monocline dip towards the Ramapo Fault in the northwest. Small faults mapped within the basalt units showed two trends, approximately due north and out². The faults are of limited lateral extent and more commonly show horizontal striae indicating strike slip movement. The strike slip movement direction was not determined because indicators were poorly developed. Similar features were observed in the Bernardsville quadrangle where the faults do not penetrate the underlying sedimentary units. Faults showed gouge zones that commonly influence ground water flow. Many other smaller scale slip zones (20-30 ft) were observed in the basalt. Most of these structures displayed horizontal movement while less common dip slip lineations also occurred. Fault movement predates surficial cover. Joint trends are shown in Figure 2.

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Description of map units

- Jp** Hook Mountain Basalt (Lower Jurassic) (Olsen, 1960a) - Dark-greenish-gray to black, generally fine-grained and very locally medium- to coarse-grained, amygdaloidal basalt composed of plagioclase, clinopyroxene, and iron-titanium oxides. Contains small spherules of tubular gas-escape vesicles, some filled by zirconite minerals or calcite, typically above flow contacts. Unit consists of at least two, and possibly as many as three major flows. Base of lowest flow is intensely vesiculated. Tops of flows are weathered and vesiculated. Maximum thickness is about 361 ft.
- Jt** Towaco Formation (Lower Jurassic) (Olsen, 1980a) - Unit is covered by surficial deposits but identified through well logs (Stanford, 1991). Elsewhere unit is reddish-brown to brownish-purple, fine to medium-grained, fine-grained sandstone, siltstone, and calcareous siltstone, and black microlaminated calcareous siltstone and mudstone. Irregular mudcracks, symmetrical ripple marks present. Sandstone is often hummocky to trough cross-laminated and siltstone commonly planar laminated or bioturbated, indistinctly laminated to massive. Several inches of unit have been thermally metamorphosed along contact with Hook Mountain Basalt. Maximum thickness is about 1,250 ft.
- Jp** Preakness Basalt (Lower Jurassic) (Olsen, 1980a) - Unit is covered by surficial deposits but identified through well logs (Stanford, 1991). Elsewhere unit is reddish-brown to brownish-purple, fine-grained, dense, hard basalt composed mostly of interlocking calcic plagioclase and clinopyroxene. Contains small spherules of tubular gas-escape vesicles, some filled by zirconite minerals or calcite, just above scoriaceous flow contacts. Unit consists of at least three major flows, the tops of which are marked by prominent vesiculated zones up to 8 ft thick. Radiating slender columns 2 to 24 in. wide, due to shrinkage during cooling are abundant near the base of the lowest flow and well developed along Interstate Route 78. A thin, 6 to 25 ft-thick bed of reddish-brown siltstone separating the first two flows has been identified elsewhere but not within this mapped area. Maximum thickness of unit is about 1,040 ft.
- Jp** Fallville Formation (Lower Jurassic) (Olsen, 1980a) - Unit is covered by surficial deposits but identified through well logs (Stanford, 1991). Elsewhere unit is reddish-brown to light grayish-red, fine to coarse-grained sandstone, siltstone, and calcareous limestone. Upper part of unit is predominantly thin- to medium-bedded, reddish-brown siltstone. Reddish-brown sandstone and siltstone are moderately well sorted, commonly cross-laminated, and interbedded with reddish-brown, planar-laminated siltstone and mudstone. Two thin, laterally continuous sequences (Jp), each up to 10 ft thick of dark-gray to black, carbonaceous limestone, light gray limestone, and medium-gray calcareous siltstone, and gray or olive, desiccated shales to silty shales occur near the base. Several inches of unit have been thermally metamorphosed along contact with Preakness Basalt (Jp). Thickness ranges from 450 to 483 ft.
- Jp** Orange Mountain Basalt (Lower Jurassic) (Olsen, 1980a) - Dark-greenish-gray to black, fine-grained, dense, hard basalt composed mostly of calcic plagioclase and clinopyroxene. Locally contains small spherules of tubular gas-escape vesicles, some filled by zirconite minerals or calcite, typically above base of flow contact. Unit consists of three major flows that are separated in places by a weathered zone, a bed of thin reddish-brown siltstone, or by volcanoclastic rock. Lower part of upper flow is locally pillowed; upper part has scoriaceous flow structures. Middle flow is massive to columnar jointed. Lower flow is generally massive with widely spaced curvilinear joints and is pillowed near the top. Individual flow contacts characterized by vesiculated zones up to 8 ft thick. Thickest of unit is about 415 ft.
- JTp** Passaic Formation (Lower Jurassic and Upper Triassic) (Olsen, 1980) - Interbedded sequences of reddish-brown, and less often maroon or purple, fine to coarse-grained sandstone, siltstone, shaly siltstone, silty mudstone, and mudstone. Reddish-brown sandstone and siltstone are thin to medium bedded, planar to cross bedded, micaceous, and locally mudcracked and ripple cross-laminated. Root casts and load casts are common. Shaly siltstone, silty mudstone, and mudstone are fine-grained, very thin to thin bedded, planar to ripple cross-laminated, locally fissile, bioturbated, and contain evaporite minerals. They form rhythmically fining-upward sequences as much as 15 ft thick. Several inches of unit have been thermally metamorphosed along contact with Orange Mountain Basalt (Jp). Unit is barely exposed in southwestern part of the map area, but regionally is as much as 11,480 ft thick.

Explanation of map symbols

- Contact - Queried where uncertain
- Contours of bedrock surface elevation, 50 foot interval
- Faults - Queried where uncertain
- High-angle fault of uncertain movement
- Strike and dip of inclined beds
- Other features
- Abandoned rock quarry
- Well numbers and locations used to define bedrock geology from Stanford (1991). See Table 1.

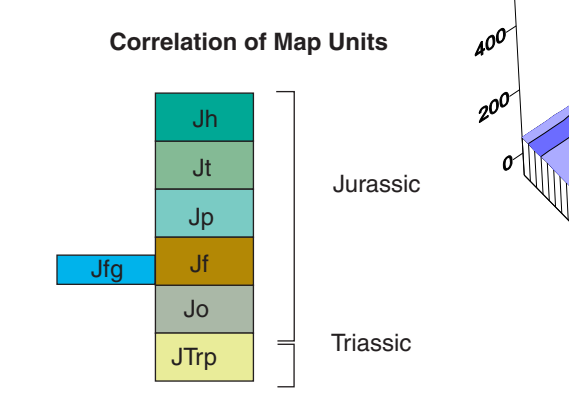


Table 1. Well information from Stanford (1991)

Well No.	Permit No.	Depth (ft)	Diater's Log Description	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100	101	102	103	104	105	106	107	108	109	110	111	112	113	114	115	116	117	118	119	120	121	122	123	124	125	126	127	128	129	130	131	132	133	134	135	136	137	138	139	140	141	142	143	144	145	146	147	148	149	150	151	152	153	154	155	156	157	158	159	160	161	162	163	164	165	166	167	168	169	170	171	172	173	174	175	176	177	178	179	180	181	182	183	184	185	186	187	188	189	190	191	192	193	194	195	196	197	198	199	200	201	202	203	204	205	206	207	208	209	210	211	212	213	214	215	216	217	218	219	220	221
59	26-6863	0-80	surficial red shale	123	25-13808	0-50	surficial red shale	187	0-23	23-271	surficial red shale	188	0-28	28-305	surficial red shale	189	0-46	46-505	surficial red shale	190	0-78	78-100	surficial red shale	191	0-8	8-57	57-80	surficial red shale	192	0-56	56-100	surficial red shale and sandstone	193	0-63	63-100	surficial red shale and sandstone	194	0-69	69-100	surficial red shale and sandstone	195	0-70	70-100	surficial red shale and sandstone	196	0-70	70-100	surficial red shale and sandstone	197	0-70	70-100	surficial red shale and sandstone	198	0-70	70-100	surficial red shale and sandstone	199	0-70	70-100	surficial red shale and sandstone	200	0-70	70-100	surficial red shale and sandstone	201	0-70	70-100	surficial red shale and sandstone	202	0-70	70-100	surficial red shale and sandstone	203	0-70	70-100	surficial red shale and sandstone	204	0-70	70-100	surficial red shale and sandstone	205	0-70	70-100	surficial red shale and sandstone	206	0-70	70-100	surficial red shale and sandstone	207	0-70	70-100	surficial red shale and sandstone	208	0-70	70-100	surficial red shale and sandstone	209	0-70	70-100	surficial red shale and sandstone	210	0-70	70-100	surficial red shale and sandstone	211	0-70	70-100	surficial red shale and sandstone	212	0-70	70-100	surficial red shale and sandstone	213	0-70	70-100	surficial red shale and sandstone	214	0-70	70-100	surficial red shale and sandstone	215	0-70	70-100	surficial red shale and sandstone	216	0-70	70-100	surficial red shale and sandstone	217	0-70	70-100	surficial red shale and sandstone	218	0-70	70-100	surficial red shale and sandstone	219	0-70	70-100	surficial red shale and sandstone	220	0-70	70-100	surficial red shale and sandstone	221	0-70	70-100	surficial red shale and sandstone																																																																												

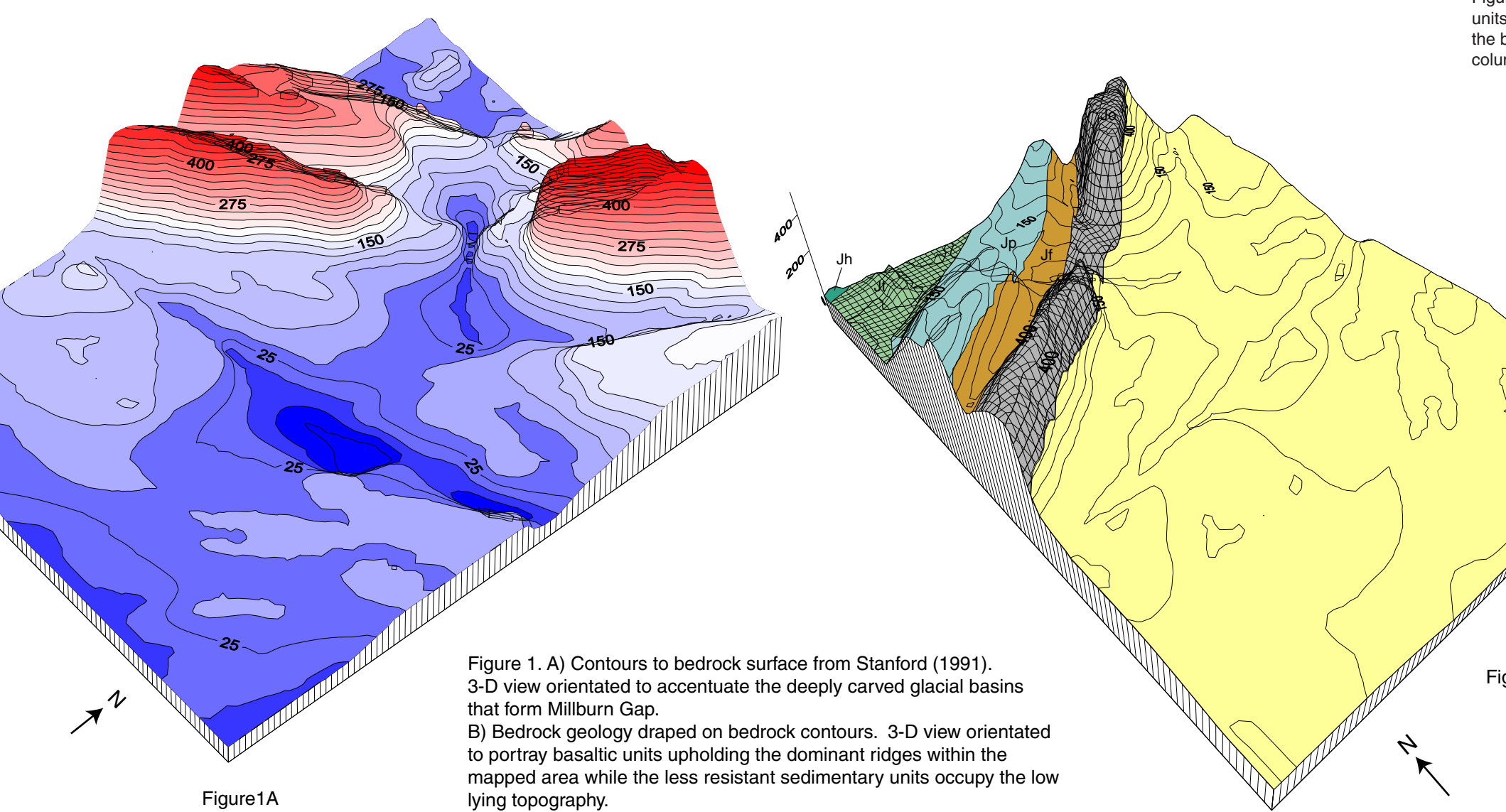
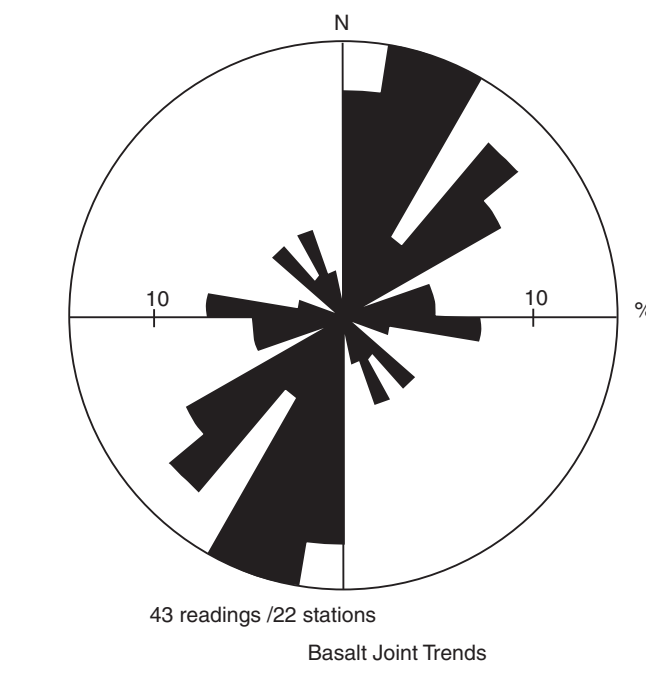
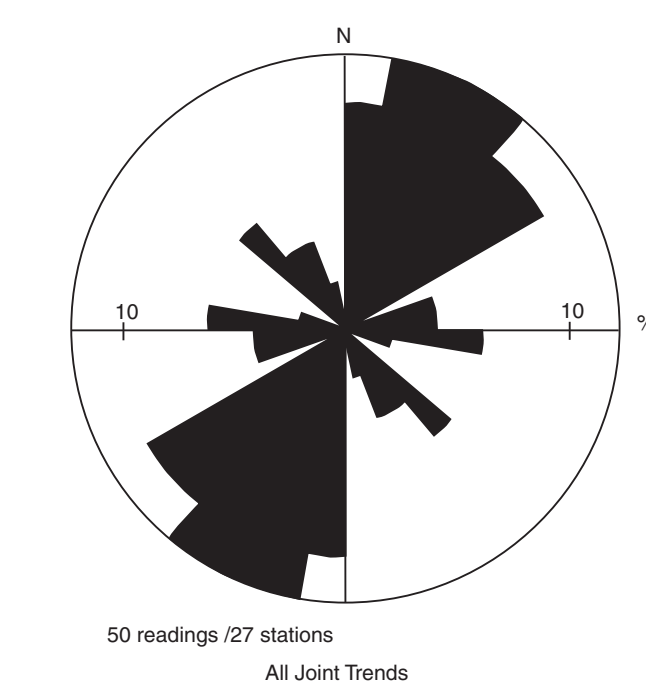


Figure 1. A) Contours to bedrock surface from Stanford (1991). 3-D view orientated to accentuate the deeply carved glacial basins that form Millburn Gap. B) Bedrock geology draped on bedrock contours. 3-D view orientated to portray basaltic units upholding the dominant ridges within the mapped area while the less resistant sedimentary units occupy the low lying topography.

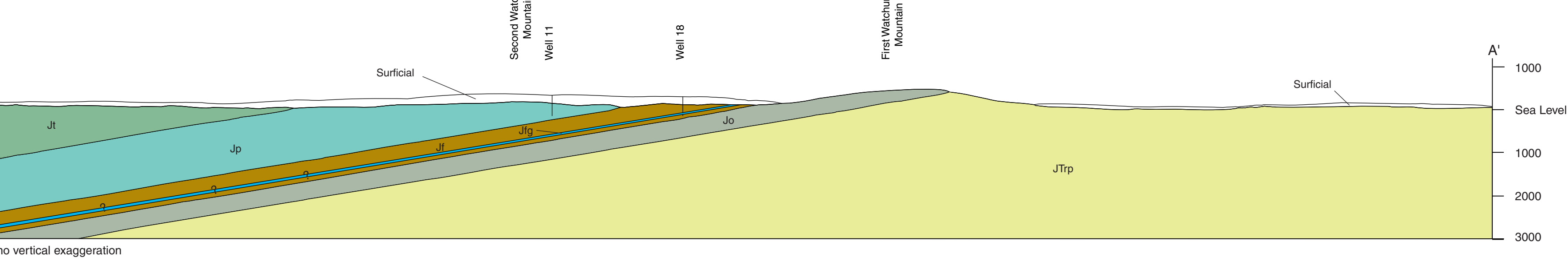


Figure 1B. Bedrock geology draped on bedrock contours. 3-D view orientated to portray basaltic units upholding the dominant ridges within the mapped area while the less resistant sedimentary units occupy the low lying topography.